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## GEOTECHNICAL PARAMETERS STUDY USING SEISMIC REFRACTION TOMOGRAPHY

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#### Abstract

In this study, correlation is made between seismic P-wave velocities (V<sub>p</sub>) with standard penetration test (SPT-N) values to produce soil parameter estimation for engineering site applications. A seismic refraction tomography (SRT) line of 69 m length was spread across two boreholes with 3 m geophones spacing. The acquired data were processed using Firstpix, SeisOpt2D and surfer8 software. The V<sub>p</sub> at particular depths were pinpointed and correlated with geotechnical parameters (SPT-N values) from the borehole records. The correlation between V<sub>p</sub> and SPT-N values has been established. For cohesive soils, it is grouped into three categories according to consistencies; stiff, very stiff and hard, having velocity ranges of 575-314 m/s, 808-1483 m/s and 1735-2974 m/s, respectively. For non-cohesive soils, it is also divided into three categories based on the denseness as loose, medium dense and dense with V<sub>p</sub> ranges of 528-622 m/s, 900-2846 m/s and 2876-2951 m/s, respectively.

Keywords: Correlation; soil strength; P-wave velocity

### Abstrak

Dalam kajian ini, halaju gelombang seismik P, (V<sub>P</sub>) dan nilai ujian penusukan piawai (SPT-N) dilakukan untuk menghasilkan penganggaran parameter tanah bagi aplikasi tapak kejuruteraan. Satu garis tinjaun seismik pembiasan tomografi yang panjangnya 69 m direntangkan merentasi dua lubang jara dengan sela antara geofon ialah 3 m. Data yang diperolehi diproses menggunakan perisian Firstpix, SeisOpt2D dan surfer8. V<sub>P</sub> pada kedalaman tertentu dikeluarkan dan dikorelasikan dengan parameter geoteknikal (SPT-N) daripada rekod lubang jara. Korelasi antara V<sub>P</sub> dan nilai SPT-N telah dihasilkan berdasarkan kekuatan tanah. Untuk tanah jeleket, ia telah dikumpulkan kepada tiga kategori berdasarkan kepada kekonsistenan; kaku, sangat kaku dan keras, masing-masing dengan halaju 575-314 m/s, 808-1483 m/s and 1735-2974 m/s. Untuk tanah tidak jeleket, ia juga dibahagikan kepada tiga kategori berdasarkan kepadatar; longgar, sederhana padat dan padat, masing-masing dengan nilai halaju 528-622 m/s, 900-2846 m/s and 2876-2951 m/s.

Kata kunci: Korelasi; kekuatan tanah; halaju gelombang-P

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#### **Full Paper**

#### **1.0 INTRODUCTION**

Geophysical methods have been widely utilized in engineering fields for subsurface study. In particular, seismic refraction (SR) method had been widely used to encounter civil engineering problems in Malaysia geotechnical as in pavement, such and environmental engineering [1]. Seismic refraction tomography (SRT) is an advanced of SR method used in interpreting seismic refraction data, which apply gridding and technique of inversion to determine velocity in 2-dimension blocks (pixels) for velocities modelling. As a result, SRT could provide better resolution and more accurately model the complex subsurface velocity of structure. The combination of geophysical and geotechnical approach was proved as a very cost effective methods [2].

Detailed information about mineral content of the soil or rock, and also the rock fabric are usually obtained through drilling. The drilling method relies on the aim of the study, information required and the soil or rock types being drilled. From the past decades, the interest for drilling through geologic structures of the uppermost continental crust has been increased [3]. Several studies have been done to acquire a relation between seismic wave velocities and soil mechanics parameters [4-11]. Also, many studies have been carried out to obtain a relation between the seismic wave velocities and various lithological properties of rocks for the exploration geophysics purpose [12-18]. Unfortunately, there is no theoretical relationship which can be used to correlate between destructive methods (e.g., SPT-N) and non-destructive methods (e.g., seismic refraction methods) [19]. The purpose of this study is to develop a correlation between seismic P- wave velocities (Vp) obtained from SRT with respect to SPT-N values.

#### 2.0 METHODOLOGY

Universiti Sains Malaysia (USM) is an investigation area, situated at east of Penang Island, Malaysia. Geologically, igneous rock underlain major portion of the Penang Island (Figure 1). The igneous rock is granites in terms of Streckeisen classification and classified on the basis of proportions of alkali feldspar to total feldspars. Granites of Penang Island are further divided into two main groups: North Penang Pluton and South Penang Pluton. In the northern part of the island, the alkali feldspars, that generally do not exhibit distinct cross-hatched twining, are orthoclase to intermediate microcline in composition. In the southern region, they generally exhibit well-developed crosshatched twining and are believed to be microcline. The North Penang Pluton has been divided into Feringgi Granite, Tanjung Bungah Granite and Muka Head micro granite. The South Penang Pluton has been divided into Batu Maung Granite and Sungai Ara Granite [20].



Figure 1 Geological map of Penang Island [20]

Seismic data were acquired on a seismic line using a 24 channels ABEM Terraloc MK8 seismograph with a, 24 geophones (14 Hz), 2 seismic cables, 20 kg weightdrop and a roll of trigger cable. In this study, geophones have been set at 3 m interval with 5 inline shot-points and 8 offset shot- points (Figure 2). The data were processed using FIRSTPIX v4.21, OPTIM (SeisOpt2D) and SURFER 8 software. Two boreholes (BH1 and BH2) are located along the seismic line at the distance of 23.3 m and 45.0 m, respectively. The correlations between  $V_p$  and SPT-N values were established for two soil categories; cohesive and noncohesive.



Figure 2 Schematic diagrams of shot points for seismic survey

#### **3.0 RESULTS AND DISCUSSION**

SRT is utilised to produce a seismic cross section which is displayed using Surfer 8 software after being processed using FirstPix and SeisOp@t2D software. Seismic results show that the  $V_p$  of the subsurface varies from 400-3600 m/s to the depth of 38 m (Figure 3). Table 1 shows borehole record of BH1 and BH2 including SPT-N value and soil description of the sample at interval of 1.50 m.





	BH 1 BH 2			BH 2	
Depth	Description	SPT N-value	Depth	Description	SPT N-value
1.50	Loose, silty GRAVEL	6	1.50	No Recovery	6
3.00	Very stiff, sandy SILT	19	3.00	Stiff, sandy SILT	9
4.50	Stiff, sandy SILT	13	4.50	Loose, SAND	8
6.00	Stiff, SILT	15	6.00	Stiff, sandy SILT	10
7.50	Very stiff, SILT	17	7.50	Stiff, sandy SILT	9
9.00	Very stiff, SILT.	14	9.00	Stiff, sandy SILT	11
10.50	Stiff, sandy SILT	19	10.50	Stiff, sandy SILT	14
12.00	Very stiff, sandy SILT	22	12.00	Very stiff, sandy SILT	16
13.50	Very stiff, sandy SILT	23	13.50	Very stiff, sandy SILT	17
15.00	Medium dense, SAND	21	15.00	Very stiff, sandy SILT	18
16.50	Medium dense, SAND.	9	16.50	Stiff, sandy SILT	15
18.00	Medium dense, SAND.	16	18.00	Medium dense , silty SAND	18
19.50	Medium dense, SAND	15	19.50	Very stiff, sandy SILT	20
21.00	No Recovery	39	21.00	Very stiff, sandy SILT	21
22.50	Very stiff, sandy SILT	15	22.50	Very stiff, sandy SILT	18
24.00	Very stiff, sandy SILT	43	24.00	Very stiff, sandy SILT	17
25.50	Stiff, sandy SILT	50	25.50	Medium dense, silty SAND	20
27.00	Hard, sandy SILT	30	27.00	Medium dense, silty / clayey SAND	23
28.50	Hard, sandy SILT	45	28.50	Very stiff, sandy SILT of	20
30.00	Stiff, sandy SILT	45	30.00	Medium dense, silty SAND of	26
31.50	Hard, sandy SILT	50	31.50	Medium dense, silty SAND	50
33.00	Hard, sandy SILT	50			
34.50	No Recovery	50			

Figure 4 shows the subsurface velocity distribution of SRT integrated with  $V_p$  and SPT-N value for BH1 and BH2 together with soil lithology at interval of 1.5 m. A graph of  $V_p$  and SPT-N value against depth for BH1 and BH2 were plotted (Figure 5). The result show direct correlation between  $V_p$  and SPT-N value in which  $V_p$  increases with SPT-N values. The lithology recorded

from BH 1 shows a complex subsurface, which may affect the SPT-N values. This has been shown by fluctuated SPT-N values with increasing depth compared to BH 2, in which the SPT-N values consistently increasing with depth since no variation in lithology occurred.



Figure 4 Subsurface velocity distribution of SRT, Vp and SPT-N value for BH1 and BH2 together with soil lithology at interval of 1.5 m



Figure 5 Relationship between  $V_p$  and SPT-N values with depth for BH 1 and BH 2

The V<sub>p</sub> and SPT-N values were group based on soil strength obtained from standard penetration test and tabulated into cohesive and non-cohesive soils group (Table 2). Cohesive soil, was grouped into three categories based on the consistencies (stiff, very stiff and hard) with V<sub>p</sub> value of 575-1314 m/s, 808-1483 m/s and 1735-2974 m/s; and SPT-N value are 9-15, 16-19 and 39-50, respectively. For non-cohesive soils, it is also divided into three categories based on the denseness

or state of packing (loose, medium dense and dense) with V<sub>p</sub> value of 528-622 m/s, 900-2846 m/s and 2876-2951 m/s; and SPT-N value are 6-8, 18-26 and 45-50, respectively. This study provides a cost effective alternative to generate a good estimation of SPT-N value using V<sub>p</sub> obtained that can be applied on other sites with similar lithology, without in situ test.

	COHESIVE						
V <sub>p</sub> (m/s)	SPT N-value	Consistency	Description				
575 1314	9 15	Stiff	sandy SILT and SILT				
808 1483	16 19	Very stiff	sandy SILT and SILT				
1735 2974	39 50	Hard	sandy SILT and SILT				
	Ν	ON COHESIVE					
V <sub>p</sub> (m/s)	SPT N-value	Denseness	Description				
528 622	6 8	Loose	silty GRAVEL and SAND				
900 2846	18 26	Medium dense	Silty SAND and SAND				
2876 2951	45 50	Dense	Silty SAND				

#### 4.0 CONCLUSION

This study proposed a cost and time effective method of providing good estimation of SPT-N values using  $V_p$ . The results showed direct proportional of SPT-N and  $V_p$  values.

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